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Keywords: *vegetation restoration; recovered chrono- sequence; soil microbial parameters; soil microbial metabolic; soil aggregate size.*

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The Distribution of Soil Microbial Parameters based on Aggregate Fractions in Successional Grassland Restoration Ecosystems on the Loess Plateau

Zhijing Xue ^α, Zhongming Wen ^σ, Zhengchao Zhou ^ρ & Shaoshan An ^ω

Abstract- Soil microbial biomass and microbial metabolic quotients have been identified as the most effective indicators of a response to ecosystem development and disturbance. In order to determine the effects of grassland restoration programs, we analyzed the distribution of soil microbial parameters in different aggregate fractions at 1yr, 10 yr. and 30 yr. Results show that C_{mic} and P_{mic} increased with grassland succession, the greatest values were concentrated at a micro-aggregate size. However, qCO_2 decreased as the years under restoration increased. qCO_2 was significantly correlated with SOC, TN, P_{mic} (negatively) and C_{mic} (positively). Natural succession of grassland alters soil microbial properties positively, which plays a major role in aggregate formation. The time under vegetation rehabilitation impacts soil microbial parameters at different aggregate sizes, especially for qCO_2 where 2-3 mm was found to be the aggregate size that was least influenced by soil quality.

Keywords: vegetation restoration; recovered chronosequence; soil microbial parameters; soil microbial metabolic; soil aggregate size.

1. INTRODUCTION

The Loess plateau of China is an erosion-prone region that is susceptible to the forces of wind and water, and has been called the “the most highly erodible soil on Earth” (Tian and Huang, 2000). The area covers 62.38×10^6 ha which includes 20.15×10^6 ha of grassland, with 91.2 % of that area being occupied by natural grassland. Grassland is the most common terrestrial ecosystem on Earth with a crucial function of regulating climate and maintaining a balanced ecosystem (Belsky, 1992). Widespread stress practice, intensive cultivation, overgrazing and large-scale monocultures have caused the Loess plateau to acquiesce to soil erosion and a series of related eco-

environment problems (Montalvo et al., 1997; Fu, et al., 2000; An et al., 2009). Environmental restoration, which aims to restore disturbed ecosystems, has been an important tool for mitigating human pressure on natural environments (Holl, et al., 2003) and for improving ecological services (Doren, et al., 2009). The Grain to Green program, a national ecological restoration program implemented in 1999, has made remarkable advances in vegetation recovery on the Loess plateau (Feng, et al., 2013). Vegetation restoration is the most effective method for abating soil erosion and degradation (Hou, et al., 2002; Montalvo, et al., 1997). Several countermeasures have been implemented including eco-environment rehabilitation using engineering and biological approaches (Wang, 2002). Natural grassland protection and restoration is one of the important parts of a vegetation restoration program (Kerri, et al., 2002).

Soil structure, especially the spatial distribution of OM within the organic-mineral matrix of soil, exhibits control over microbial mediated decomposition processes in terrestrial ecosystems (Oades, 1988; Van Veen and Kuikman, 1990; Golchin, 1994; Ladd, et al., 1996). It exerts a significant amount of influence over other edaphic conditions and the surrounding environment, and often imbues a degree of stability to aggregates (Bronick and Lal, 2005). Soil aggregates are structural units, where a group of primary soil particles cohere to each other more strongly than to other surrounding particles (Andraski and Scanlon, 2002). Research indicates that soil aggregates greatly impact the soil microbial biomass and mineral nutrients reserves (Hernández-Hernández and López-Hernández, 2002; Villar, et al., 2004). Thus, the soil microbial biomass demonstrates a similarly positive relationship with soil structure, and microbial biomass shows significantly positive relationships with aggregate size and stability (Gupta & Germida, 1988; Drury, et al., 1991; Jocteur-Monrozier, et al., 1991; Ghoshal and Singh, 1995). When assessing the effects of disturbances on soil quality, Anderson and Domsch (1985) proposed the ratio of soil basal respiration to microbial biomass (microbial metabolic quotient,

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specific respiration of the biomass or qCO_2), conceptually based on Odum's theory of ecosystem succession, is increasingly being used as an index of ecosystem development (Odum, 1969 and 1985; Wardle & Ghani, 1995). Soil respiration is a major determinant of the carbon balance, Hunt, et al. (2004) and Wohlfahrt, et al. (2005) show that the proportionally largest emission of CO_2 comes from grassland soil. Furthermore, soil microbial biomass basal respiration and microbial metabolic quotient all respond readily to disturbance effects and provide an effective early warning for the deterioration of soil quality (Powelson, et al., 1987; Wardle, 1992; Wardle & Ghani, 1995). They have been used as indicators of change in soil organic matter that occur in response to land use alteration, tillage practices and soil pollution (Sparling et al., 2003; An et al., 2009).

We hypothesize that the years under grassland restoration effects soil structure formation which consequently effects the distribution of soil microbial biomass and microbial metabolic quotient. The objectives of this study are (1) to investigate the distribution of soil microbial biomass and microbial

b) Soil sampling collection

Samples were collected from 9 sites during May 2016 in each of the three restoration sites. For each restoration site (characterized by year), three samples/location were collected as replicates. An area of 50 m×50 m was selected for each of the 9 sites; within this area, three 10 m×10 m plots were selected and S-type multiple sampling methods were used for soil collection (0-20 cm). Soil sample collections were repeated 5 times in the field, mixed, transported to the laboratory and sieved (>5, 5-3, 3-2, 2-1, 1-0.25 and <0.25 mm). Care was taken during sample preparation to minimize moisture loss and samples were divided in two. One section was preserve in 4°C for soil microbial measurements. The other was air dried for basic soil characteristics analysis. The water-holding capacity was determined by saturating each 100 cm³ turf sample with water and allowing it to drain at field capacity under cover for 48 h at ambient temperature; Duplicate soil

metabolic quotient in different soil aggregate hierarchies at different restoration periods; (2) to determine the correlation between basic soil characteristics, and soil microbiological parameters; and (3) in this context, to examine soil aggregate stability.

II. MATERIALS AND METHODS

a) Study site description

The study area was located in Chinese Loess Plateau, the south of Ningxia Province (106° 25'--106° 30'E, 35° 59'--36° 20'N). The region has a sub-arid climate characterized by seasonal rainfall with periodic local flooding and drought; the average annual temperature was 6°C, and the average annual rainfall was 400-450 mm. The rainy season lasted from July to September and the rainfall in July accounting for 24 % of the annual rainfall. Most of the land is at altitude of 1800-2040 m and is closely dissected by steep galleys (An et al., 2009). Grassland soils from three restoration years (30 years, natural grassland; 10 years, natural grassland; 1 year, abandoned grassland), were processed at the Yunwu Observatory for Vegetation Protection and Eco-environment.

Table 1: Sampling site characteristics

Successional Years	Dominant Species	Accompanying species	Elevation (m)	Slope gradient (°)
30	Thymus mongolicus Stipa bungana	Stipa grandiss Artemisia sacrorum. Artamisia scoparia	1930	12
10	Leymus secalinus Thymus mongolicus	Artemisia scoparia Potentilla bifurca Stipa bungana	1908	11
1	Artemisia scoparia Potentilla bifurca	Thymus mongolicus Heteropappus altaicus Thymus mongolicus Stipa bungana	1940	10

cores were dried at 105°C overnight to determine moisture content (An et al., 2009).

c) Sampling analysis

The soil samples were air dried and passed through a 2 mm sieve. The basic soil characteristics were analyzed through soil chemical and physical analyses (ISSCAS, 1981). Soil organic carbon was measured by wet digestion in a mixture of 5 ml of 0.136 mol/L potassium dichromate and 5 ml of concentrated sulfuric acid. Soil total nitrogen was measured using Kjeldahl digestion. Alkali-ExTR-N was measured using a micro-diffusion method in which NH₃ was released from the soil sample by NaOH and then absorbed by boric acid. The ammonium borate product was titrated with 0.01 M HCl. Available phosphorus (Av-P) was extracted and measured in a buffered alkaline solution containing 0.5 M sodium bicarbonate. The extracts were quantified calorimetrically at 660 nm with a spectrophotometer

(UV2300, Hitachi, Tokyo, Japan) (Xue et al., 2014). The readily available potassium was extracted from the soil with 1 mol/L NH₄OAc and was measured using flame photometry. The soil pH was measured in water (1/2.5 w/v), and the moisture content was assessed by drying the samples in an oven at 105°C overnight.

Soil basal respiration was estimated through CO₂ evolution at 25 °C in samples incubated for 10 days (Jenkinson and Powlson, 1976a). Measurements were made in the laboratory at 50 % of field water holding capacity (WHC). CO₂ respired was trapped in NaOH, and the residual NaOH was titrated with HCl (An et al., 2009).

The soil microbial biomass C, N and P levels were determined using the fumigation-extraction method. A 15 g sample of oven-dried, field-moist-equivalent soil (<2 mm) was fumigated with chloroform for 24 h with the chloroform being removed by repeated evacuation. The soil was reinoculated with a small amount of unfumigated soil and incubated at a constant temperature (usually 22 or 25) for 10 days at field capacity or at 50 % of its water-holding capacity (approximately -0.01 MPa). An additional, unfumigated soil sample was used as a control (Jenkinson, 1976; Jenkinson and Powlson, 1976b; Xue et al., 2014). The soil microbial biomass carbon level was determined

immediately using a TOC analyzer or the sample was preserved at -18 (Phoenix 8000, Tekmar Dohrmann, Mason, OH, U.S.A.). A k_C value of 0.4 was chosen to calculate the total soil microbial biomass C (Wu et al., 1990). To determine soil microbial biomass N, fumigated and unfumigated samples were extracted via potassium persulfate oxidation and then measured using ultraviolet spectroscopy. The extract liquor was digested with alkaline auto-oxidation using 0.15 mol/L NaOH and 30 g/L K₂S₂O₈ with a 1:1 v/v mixture of oxidant and soil at 120-124 for 30 min (Zhou and Li, 1998; Xue et al., 2014). The concentration of N was determined colorimetrically using a spectrophotometer (UV2300, Hitachi, Tokyo, Japan) at 220 and 275 nm. A k_N value of 0.54 was chosen to calculate the total soil microbial biomass N (Vance et al., 1987). The soil microbial P level was determined calorimetrically with a spectrophotometer. 2.5 g of fumigated and unfumigated soil was placed into a 150-ml flask containing 50 ml of 0.5 mol/L NaHCO₃ solution and 2 g of P-free active charcoal. 5 ml of Mo-Sb spectrochrometry solution was added for color development. After 30 minutes, the color was determined with a spectrophotometer (UV2300, Hitachi, Tokyo, Japan) at 700 nm. The microbial P was calculated using a k_P factor of 0.40 (Hedley and Stewart, 1982; An et al., 2012).

III. RESULTS

a) Basic soil characteristics

Table 2: Basic soil characteristics

Soil Characteristics	Successional years		
	30 years	10 years	1 year
BD	0.91	1.13	1.28
MFC	37.08	31.83	27.34
Por.	54.85	54.80	50.36
SOC	27.85	18.15	9.50
TN	3.10	1.88	1.17
Ap-P	3.56	3.37	2.71
Ap-K	166.24	241.69	153.04
NH ₄ ⁺ -N	14.08	4.21	4.53
NO ₃ ⁻ -N	16.62	6.40	15.06
C : N	8.87	8.26	8.26
Inv.	20.40	14.77	7.83
Alk.-P	12.43	9.99	5.95
Ure.	2.62	4.13	3.64
C _{mic}	1065.49	683.72	358.96
N _{mic}	25.78	41.41	17.23
P _{mic}	19.99	15.74	2.56

BD: soil bulk density (g•cm³); MFC: maximum field capacity (%); Por.: porosity (%); SOC: soil organic carbon (g•kg⁻¹); TN: total nitrogen (g•kg⁻¹); Ap-P: rapid available phosphorus (mg•kg⁻¹); Ap-K: rapid available potassium (mg•kg⁻¹); NH₄⁺-N: soil ammonium nitrogen (mg•kg⁻¹); NO₃⁻-N: soil nitrate nitrogen (mg•kg⁻¹); Inv.: invertase (mg•g⁻¹); Alk.-P: Alkal-phosphatase(mg•g⁻¹); Ure.: urease (mg•g⁻¹); C_{mic}: soil microbial biomass (mg•kg⁻¹); N_{mic}: soil microbial nitrogen (mg•kg⁻¹); P_{mic}: soil microbial phosphorus (mg•kg⁻¹).

General soil characteristics for the various restoration years are shown in Table 2. As grassland recovery years increased, maximum field capacity (MFC), porosity (Por.), soil organic carbon (SOC), total nitrogen (TN), rapid available phosphorus (Ap-P), invertase (Inv), Alkal-phosphatase (Alk.-P), soil microbial biomass carbon (C_{mic}) and phosphorus (P_{mic}) also increased. However, restoration years also resulted in a decrease in bulk density (BD). The highest

concentrations of soil microbial biomass nitrogen (N_{mic}), urease (Ure) and rapid available potassium (Ap-K) were found in 10 year grassland. The distribution of aggregate size class percentage varied among the grassland restoration sites (Fig. 1). The percentage of < 0.25 and > 5 mm aggregate sizes in 30 year natural grassland was lower than in 10 year. Lastly, soil aggregate composition in 30 year grassland was mainly concentrated between 0.25 and 5 mm.

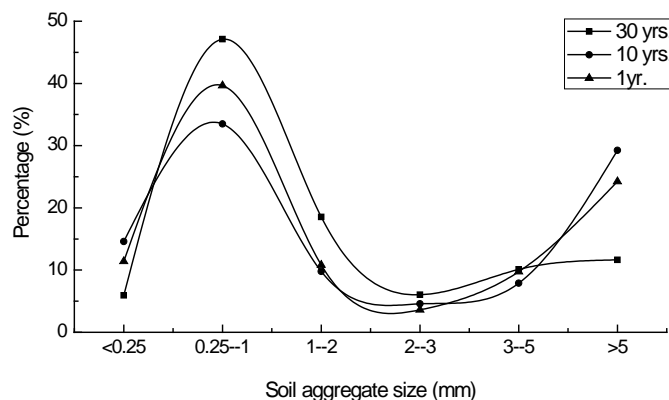
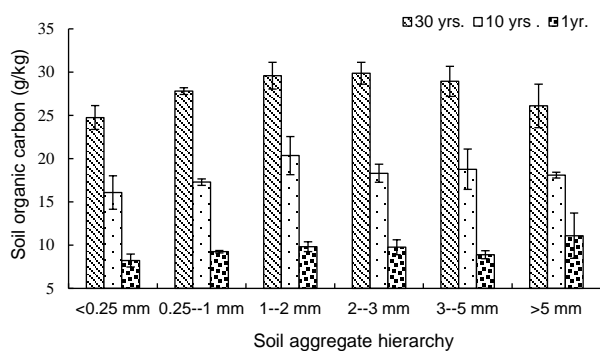


Fig. 1: Effect of restoration years on the distributions of aggregate sizes' percentage

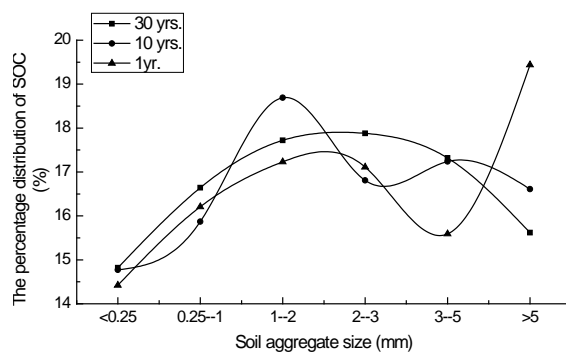
b) Distribution of SOC and TN in various aggregate sizes

The concentration of SOC and TN associated with soil grain size increased with the number of years of grassland succession (Fig. 2 (a) and (c)). For 10- and 30- year natural grassland, the percentage of SOC in 1-2, 2-3 and 3-5 mm was higher than for other aggregate sizes. In 1 year abandoned grassland, the macro-aggregate size (>5 mm) showed the greatest SOC with

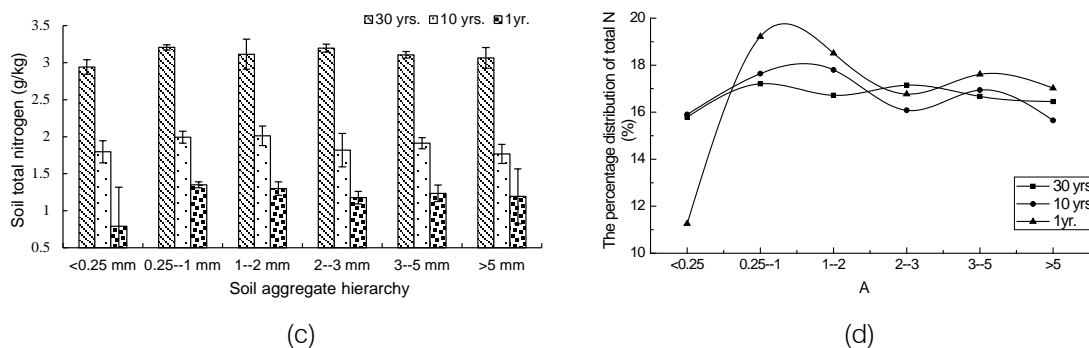
rapid enhancement between 3-5 and 5 mm (Fig. 2 (b)). The trend in distribution of TN among the various aggregate classes in 10- and 30- year natural grassland was similar (i.e., stable between 15.79 - 17.64). However, significant changes occurred in 1-year-abandoned grassland (Fig. 2 (d)). The percentage of TN (11.26) was lowest for micro-aggregates (<0.25 mm), highest in 0.25-1 mm (19.22), and remained at a lower yet steady level for 10- and 30- year (17.02-18.51).



(a)



(b)



- (a) The distribution of SOC in different vegetation restoration years;
- (b) The percent distribution of SOC in soil aggregate size
- (c) The distribution of TN in different vegetation restoration years;
- (d) The percent distribution of TN in soil aggregate size

Fig. 2: Effect of restoration years on SOC and TN for different aggregate sizes

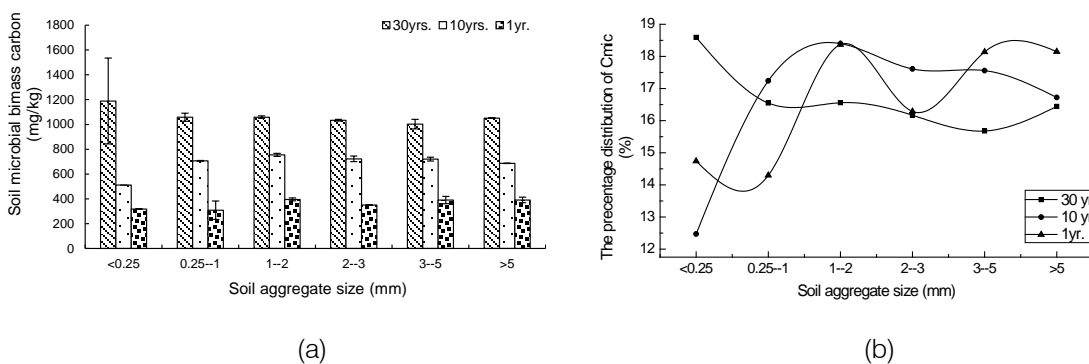
c) The distribution of soil microbial parameters at different aggregate sizes

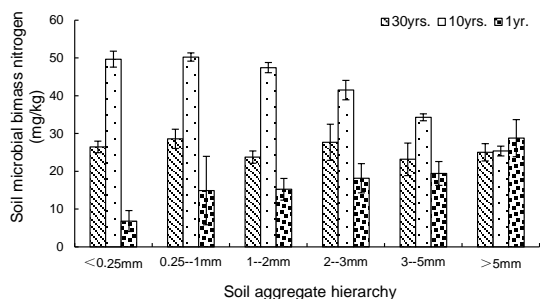
i. Soil microbial biomass

The concentrations of C_{mic} and P_{mic} at different grain sizes increases with length of grassland succession (Fig. 3 (a) and (e)). The situation with N_{mic} is different where it had the highest value in a 10 year natural grassland (Fig. 3 (c)).

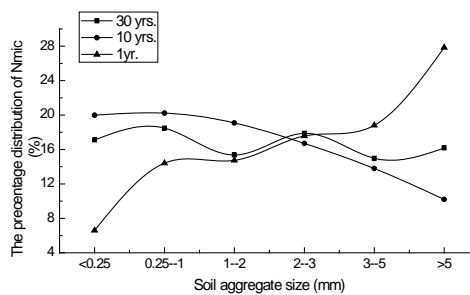
Under a 30 year natural grassland, the percentage of C_{mic} (18.59) was greatest for a micro-aggregate size of <0.25 mm, as aggregate size increased, percentages dropped to a stable range between 15.68-16.56. At 10 years, a grain size of 1-2 mm showed the greatest percentage of C_{mic} (18.4), which was a rapid increase from <0.25. The percentage of C_{mic} for other particle sizes remained at a stable level

between 16.72-17.61 (Fig 3 (b)). The curve of C_{mic} versus aggregate size for a recently abandoned grassland fluctuated greatly. There was a significant decrease in percentage of C_{mic} in 2-3 mm versus 1-2, 3-5 and >5 mm. In a 30-year natural grassland, the percentage of N_{mic} remained stable between 14.99 and 18.48 for the various aggregate sizes. As soil aggregate size increased, the distribution of N_{mic} in 1- and 10- year grassland showed a rising tendency and declining trend, respectively (Fig 3 (d)). The percentage of P_{mic} in different aggregate sizes for 10- and 30- year natural grassland remained stable. Whereas, for the highest percentage occurring in aggregate sizes of 2-3 mm, the 1-year curve showed a "normal distribution" which was significantly enhanced from 0.25-1 mm and rapidly declined at 2-3 mm (Fig 3 (f)).

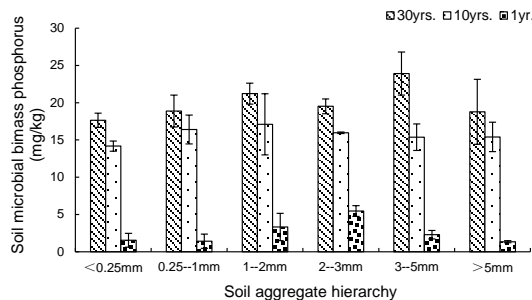




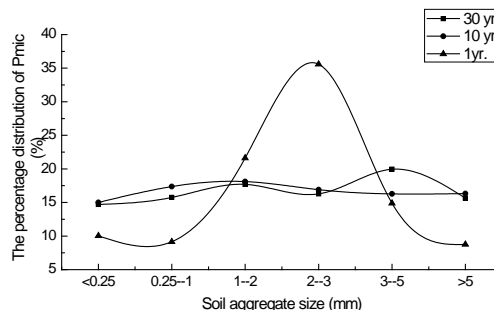
(c)



(d)



(e)



(f)

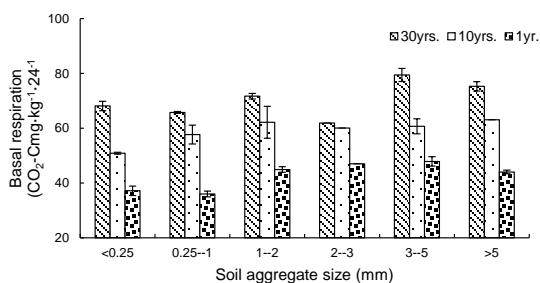
- (a) The distribution of Cmic in different grassland restoration years;
- (b) The percentage distribution of Cmic in soil aggregate size
- (c) The distribution of Nmic in different grassland restoration years;
- (d) The percentage distribution of Nmic in soil aggregate size
- (e) The distribution of Pmic in different grassland restoration years;
- (f) The percentage distribution of Pmic in soil aggregate sizes

Fig. 3: Effect of restoration years on Cmic, Nmic and Pmic in different aggregate sizes

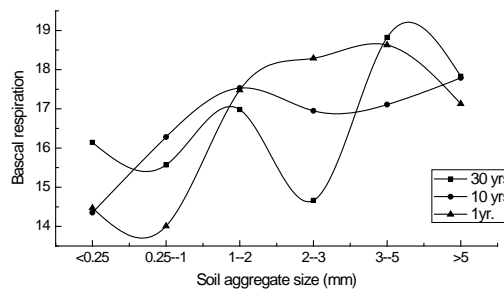
d) Soil microbial basal respiration and metabolic quotient

Soil microbial basal respiration (SBR) was enhanced as the successional years increased (Fig.4 (a)). For a grassland abandoned for 1 year, the percentage of SBR in <0.25 and 0.25-1 mm was lower than in other aggregate sizes. SBR increased with aggregate size, the highest occurring between 3-5 mm. The curve of SBR for a 30-year natural grassland for different aggregate size classes fluctuated, it was the

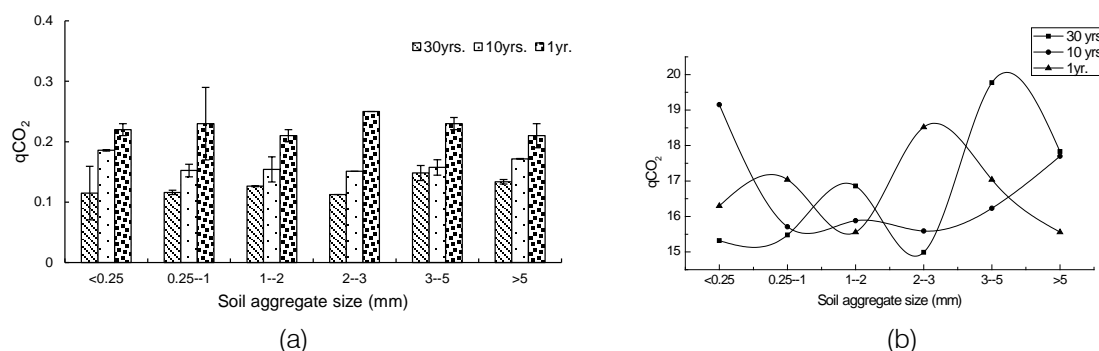
highest in 3-5 mm and lowest in 2-3 mm. However, the percentage at 10 years increased with soil aggregate size (Fig.4 (b)). Soil microbial metabolic quotient (qCO_2) among the aggregate size classes had the opposite reaction as restoration years increased (Fig.4 (c)). In a grassland abandoned for 10 years, the percentage of qCO_2 , in aggregate size classes of 0.25-1, 1-2 and 2-3 mm, was lower than in the micro-aggregate (<0.25 mm) and macro-aggregate (3-5 mm and >5 mm) classes (Fig.4 (d)).



(a)



(b)



- (a) The distribution of SBR for different grassland restoration years;
- (b) The percent distribution of SBR in soil aggregate sizes
- (c) The distribution of qCO₂ for different grassland restoration years;
- (d) The percent distribution of qCO₂ in soil aggregate sizes

Fig. 4: Effect of restoration years on SBR and qCO₂ in different aggregate sizes

e) The ratios of C/N, C_{mic}/N_{mic} and C_{mic}/P_{mic}

Table 3: The ratios of C/N, C_{mic}/N_{mic} and C_{mic}/P_{mic} for different aggregate sizes

Successional years	Size classes (mm)	C/N	C _{mic} /P _{mic}	C _{mic} /N _{mic}
30	<0.25	8.42	67.38	44.92
	0.25--1	8.67	56.05	37.03
	1--2	9.51	49.89	44.57
	2--3	9.35	52.98	37.37
	3--5	9.32	41.93	43.24
	>5	8.52	56.01	42.02
Mean		8.97	54.04	41.53
10	<0.25	8.95	36.09	10.30
	0.25--1	8.67	43.11	14.08
	1--2	10.12	44.13	15.91
	2--3	10.08	45.25	17.41
	3--5	9.81	46.87	21.02
	>5	10.23	44.53	27.04
Mean		9.64	43.33	17.62
1	<0.25	10.40	206.01	46.49
	0.25--1	6.85	218.87	20.67
	1--2	7.56	119.22	25.96
	2--3	8.29	64.21	19.32
	3--5	7.19	170.75	20.10
	>5	9.28	290.95	13.58
Mean		8.26	178.34	24.35

The ratios of C/N, C_{mic}/N_{mic} and C_{mic}/P_{mic} for the different aggregate sizes in the three successional periods were shown in Table 3. The greatest ratios of C/N, C_{mic}/N_{mic} and C_{mic}/P_{mic} were exhibited in the 10 year, 30 year and 1 year grasslands, respectively.

For the 10- and 30- year natural successional grasslands, the ratios of C/N in aggregate size of 1-2, 2-3 and 3-5 mm were higher than the others. The ratios of C_{mic}/P_{mic} in grassland under succession for 30 years

was greater in <0.25, 0.25-1 and >5 mm and were between 43.11- 46.25 in grassland under succession for 10 years, expect <0.25 mm which had the lowest value. For a grassland under restoration for one year, the ratios show a V shaped distribution among the soil aggregate sizes, with the lowest ratio occurring in 2-3 mm. The ratio of C_{mic}/N_{mic} in grassland that was in a natural state for 30 years was maintained at a stable level (37.03-42.02), but, at 10 years, the ratio increased

as soil aggregate size increased. For a grassland abandoned for 1 year, the ratio was greatest for <0.25

mm and lowest for >5 mm, the other ratios were between 9.32 - 25.96.

f) The relationship of soil microbial parameters and chemical characteristics

Table 4: Correlation coefficients between basic characteristics and soil microbial parameters

	SOC	Total N	NO ₃ -N	NH ₄ -N	Ap-K	Ap-P	C/N
C _{mic}	0.965**	0.964**	0.155	-0.269	0.087	0.574*	0.259
N _{mic}	0.268	0.206	-0.704**	-0.106	0.046	0.035	0.312
P _{mic}	0.933**	0.883**	-0.106	-0.208	0.406	0.677**	0.422
SBR	0.910**	0.882**	0.046	-0.098	0.207	0.672**	0.326
qCO ₂	-0.932**	-0.906**	0.035	0.395	-0.267	-0.473**	-0.463

*Significant at the 0.05 level; **Significant at the 0.01 level.

Significant correlations were found between soil microbial properties and chemical characteristics. As Table 4 shows, a correlation exists among the soil microbial properties that were significantly correlated with SOC, TN and Ap-P. N_{mic} did not relate to SBR and qCO₂, and only correlated with NO₃⁻-N. C_{mic} and P_{mic}

were significantly correlated with SBR, qCO₂ and C_{mic}/N_{mic}, (p<0.01). The ratio of C_{mic}/P_{mic} was negatively correlated with C_{mic} (R= -0.638, p<0.01), N_{mic} (R= -0.507, p<0.05), P_{mic} (R= -0.833, p<0.01) and SBR (R= -0.727, p<0.01), and positively correlated with qCO₂ (R=0.635, p<0.01).

Table 5: Correlation coefficients for soil microbial biomass and respiration

	C _{mic}	N _{mic}	P _{mic}	SBR	qCO ₂	C _{mic} /N _{mic}	C _{mic} /P _{mic}
C _{mic}	1						
N _{mic}	0.250	1					
P _{mic}	0.903**	0.486**	1				
SBR	0.917**	0.305	0.933**	1			
qCO ₂	0.954**	-0.414	-0.901**	0.841**	1		
C _{mic} /N _{mic}	0.593**	0.548**	0.369	0.472*	-0.463	1	
C _{mic} /P _{mic}	-0.638**	-0.507*	-0.833**	-0.727**	0.635**	0.142	1

*Significant at the 0.05 level; **Significant at the 0.01 level.

Table 5 shows that soil microbial metabolic quotient (qCO₂) was positively related to C_{mic} (R= 0.954) and negatively correlated to P_{mic} (R= -0.901), the correlation ratios were significant (p<0.01). The correlation between C_{mic} and qCO₂, P_{mic} and qCO₂ in

different aggregate sizes is shown in Figures 5 and 6. The correlation coefficients for C_{mic} and qCO₂, P_{mic} and qCO₂ in <0.25, 0.25-1, 1-2 and >5 mm were higher (0.9177-0.9804, 0.8288-0.989) than in 2-3 (0.7427, 0.7679) and 3-5mm (Fig.5 (d) and (f)).

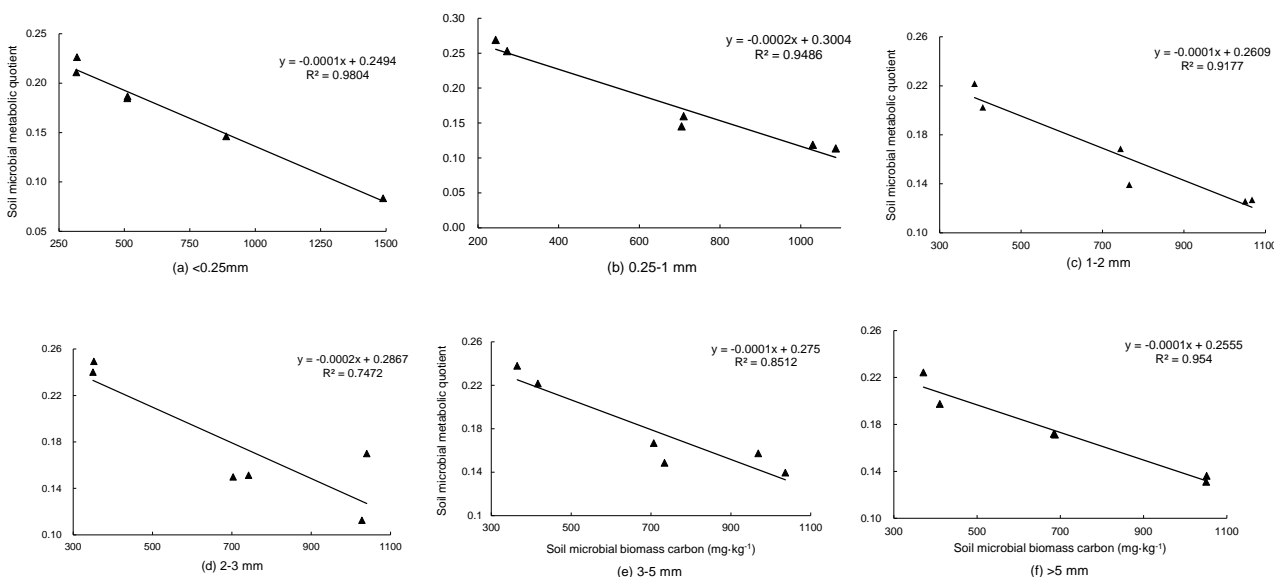


Fig. 5: The relationship between Cmic and qCO2 among different aggregate sizes

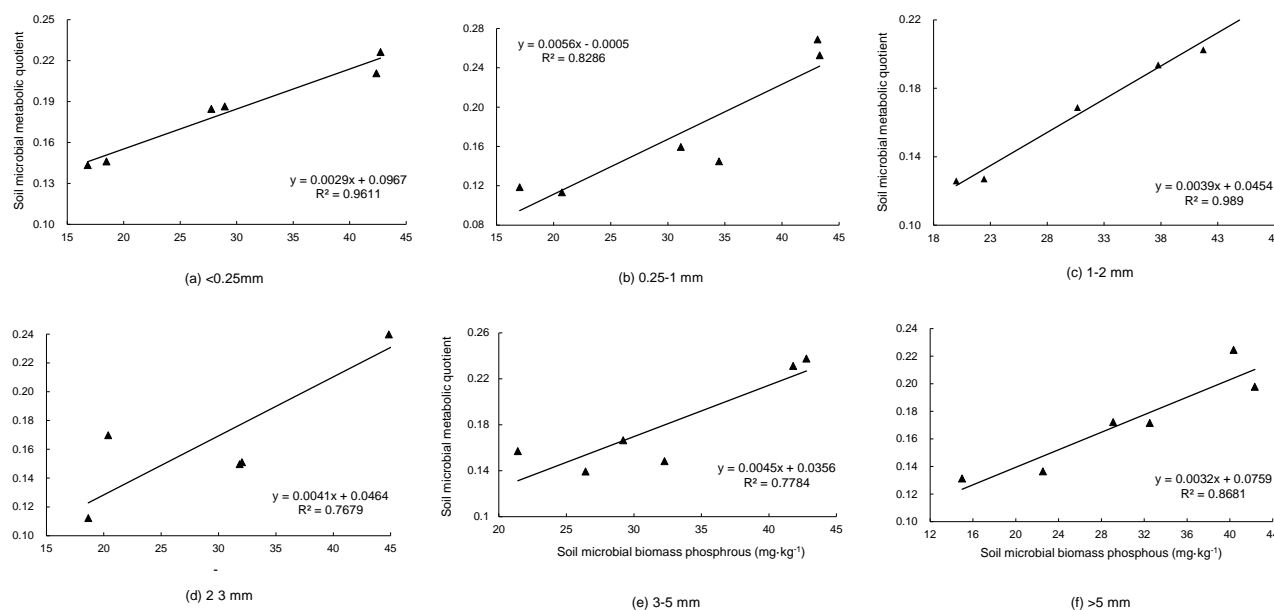


Fig. 6: The relationship between P_{mic} and qCO₂ among different aggregate sizes

IV. DISCUSSION

Soil microbial parameters are indicators of soil quality and can be used to evaluate the success of revegetation in disturbed soil ecosystems (An, et al., 2009). Soil microbial biomass, basal respiration and metabolic quotient have been previously considered the most sensitive parameters for evaluating natural and degraded systems (Chen, et al., 2000; Dick et al., 1996; Bolinder et al., 1999; An et al., 2009). The microbial parameters measured in this study were measured at different aggregate sizes, and used to quantify soil quality differences under different successional stages or lengths of revegetation. In the present study, SOC, TN and Ap-P (chemical characteristics); Inv., Alk-P (soil enzyme activities); C_{mic} and P_{mic} (soil microbial biomass) increased with an increase in restoration years. Natural grassland flourished when grazing was prohibited. After the grassland was protected with fences, plant biomass and height increased, with a concomitant increase in soil organic matter.

The arrangement of the soil aggregates and their stability had a large influence on soil properties (Lynch and Bragg, 1985). The composition of the soil aggregates influenced the amount of soil organic matter, which in return affected the soil aggregate structure and stability by functioning as a bonding agent between mineral soil particles (Haynes and Beare, 1997; Chevallier, et al., 2004; Diaz, et al., 1994). Soil organic matter protects the soil surface against raindrop impact, improves water infiltration and impacts the hydrophobic characteristics that reduce wetting rates and slaking (Angers, et al., 1998). In 10- and 30- year natural grassland, the amount of C_{mic} was greatest in < 0.25 mm, then <0.25 and lastly 1-2 mm, respectively. In

grassland that was abandoned for 1 year, the amount of C_{mic} was greatest in in 1-2, then 3-5 and, lastly >5 mm. Micro-aggregate size impacted C_{mic} accumulation, the effect of accumulation increased as the restoration years increased.

Soil microbial biomass contributed, with a significant portion in the active pools of SOC and TN. Soil microbial biomass C is one of the comprehensive indicators which reflect soil nutrient availability, biological activity and distribution of microorganism (An et al., 2009). It is the dynamic part of the soil carbon pool, which indicates the condition of soil organic matter accumulation (He, et al. 1997; Insam and Domsch, 1998). The soil microbial community influences soil mineralization and immobilization processes. The vegetative species and composition change as the vegetation restoration years increase, which impacts the soil microbial community in a direct way. An extensive root system provides the most improvement in soil structure by binding macro-aggregates to fine roots and VAM fungal hyphae and by binding micro-aggregates with adhesive bacterial metabolic products. (Lynch and Bragg, 1985; Carter, 1986; Perfect et al., 1990). Carbon input is the limiting factor in production of soil microbial biomass (Lynch and Panting, 1980; Insam &Haselwandter, 1989). The natural grassland with 30 years of restoration has a species-rich plant community that protects soil from physical disturbance (Pères, et al., 2013), and higher soil organic matter supports an abundant carbon source for the growth of microorganisms and metabolism (Six, et al., 1999 and 2000a; Anderson and Gray, 1990), which results in the highest amounts of C_{mic} and P_{mic}. The percentage of C_{mic} at different aggregate sizes was stable (Six, et al., 2000b).

Anderson and Domsch (1985) proposed that the ratio of soil basal respiration to microbial biomass (soil microbial metabolic quotient or qCO_2), was the most sensitive indicator of quantitative and qualitative changes in microbial communities caused by various management systems (Insam et al., 1989). It was identified as the most important parameter in analyzing vegetation and soil microbial characteristics (Bastida et al., 2008; An et al., 2009). Soil respiration relies on a suite of complex processes contributing to CO_2 efflux from soil surface, mainly from plant roots and microorganisms (Jia, et al., 2006; An et al., 2009). Many studies showed the changes in soil respiration and soil microbiological properties with succession (Schafer, et al., 1979; Mathes and Schriefer, 1985; Stroo and Jencks, 1982; Insam & Haselwandter, 1989). In this study, soil basal respiration was enhanced as natural restoration years increased with the highest concentrations found at the macro-aggregate level. The findings of Insam and Haselwandter (1989) on two 50-year-old soil chronosequences on reclamation sites are similar. However, soil microbial metabolic quotient (qCO_2) declined as vegetation restoration years increased. The percentage of qCO_2 in micro-aggregate sizes was lower than in macro-aggregate sizes, the distributions of SOC and TN in different aggregate sizes was similar. Because qCO_2 was negatively correlated with SOC and total N, the decrease in qCO_2 with time for the vegetated areas at different aggregate sizes resulted in competition for the available carbon source, which seemed to favor aggregate classes that needed the least amount of energy for maintenance and growth (Insam and Haselwandter, 1989). In the current study, relationships existed between qCO_2 and SOC, TN, Ap-P, C_{mic} , P_{mic} and SBR. Especially for C_{mic} and P_{mic} , the ratios in 2-3 mm were lowest (0.7472 and 0.7679, respectively). Wardle (1992) indicated that a lower microbial biomass concentration can be characteristic of disturbance, and it is an obvious factor that induces a rise in qCO_2 . A higher qCO_2 and a lower microbial biomass may both occur late in ecosystem succession, and the relationship between the two reflects their dual response to underlying stress (Wardle & Ghani, 1995). The result demonstrates that 2-3 mm is the relative stable aggregate size class in ecosystem succession.

V. CONCLUSION

Loess plateau ecosystems are suffering serious environmental problems including natural grassland degradation due to soil and water erosion. This study shows that grassland has a positive effect on soil status as indicated by an increase in the concentration of soil nutrients and microbial biomass. As grassland restoration years increased, soil microbial biomass carbon, phosphorus and basal respiration increased. Soil microbial metabolic quotient (qCO_2) an alternative

measure that changes when microbial biomass responses to disturbances, is significantly correlated to C_{mic} and P_{mic} , the lowest ratios illustrated that 2-3 mm was the stablest aggregate size in the process of ecological restoration.

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